

Lecture 8 Quasi-Geostrophic (QG) Prediction of Geopotential Tendency

8.1 Geopotential Tendency Equation

(Equation editor: $D/Dt = \partial/\partial t + u\partial/\partial x$)

Purpose: To derive a prognostic equation for predicting geopotential tendency.

➤ Based on the hydrostatic equation

$$\frac{\partial \phi}{\partial p} = -\frac{RT}{p} \quad (6.2)$$

We have

$$T = -\frac{p}{R} \frac{\partial \phi}{\partial p}$$

Substituting it into the QG thermodynamic equation

$$\frac{\partial T}{\partial t} = -u_g \frac{\partial T}{\partial x} - v_g \frac{\partial T}{\partial y} + \left(\frac{\sigma p}{R}\right) \omega + \frac{J}{c_p} \quad (6.13)$$

leads to

$$\frac{\partial \chi}{\partial p} = -V_g \cdot \nabla \frac{\partial \phi}{\partial p} - \sigma \omega - \frac{\kappa J}{p} \quad (6.22)$$

where $\kappa = R/c_p$.

- Equation (6.22) is also called “hydrostatic thermodynamic equation”.

Q: What is the physical meaning of individual terms of (6.22)?

- Equations (6.22) and the QG vorticity equation (derived in Ch.7.2 of Holton)

$$\frac{\partial \zeta_g}{\partial t} + u_g \frac{\partial \zeta_g}{\partial x} + v_g \frac{\partial \zeta_g}{\partial y} = f_0 \frac{\partial \omega}{\partial p} - \beta v_g \quad (6.18)$$

OR

$$\frac{1}{f_0} \nabla^2 \chi + u_g \frac{\partial}{\partial x} \left(\frac{1}{f_0} \nabla^2 \phi \right) + v_g \frac{\partial}{\partial y} \left(\frac{1}{f_0} \nabla^2 \phi \right) = f_0 \frac{\partial \omega}{\partial p} - \beta v_g \quad (6.18)'$$

form a closed set of equations of ϕ and ω since

$$\zeta_g = \frac{1}{f_0} \nabla_p^2 \phi \quad (6.13)$$

- Eliminate $\omega \Rightarrow$ geopotential tendency (χ) equation
 - \Rightarrow To predict geopotential height tendency
- Eliminate $\chi \Rightarrow$ Omega (ω) equation
 - \Rightarrow To diagnose vertical motion

The geopotential tendency equation can then be derived

$$\left[\nabla^2 + \frac{\partial}{\partial p} \left(\frac{f_o^2}{\sigma} \frac{\partial}{\partial p} \right) \right] \chi = -f_o V_g \cdot \nabla \left(\frac{1}{f_o} \nabla^2 \phi + f \right) - \frac{\partial}{\partial p} \left[\frac{-f_o^2}{\sigma} V_g \cdot \nabla \left(-\frac{\partial \phi}{\partial p} \right) \right] - (f_o^2 \kappa) \frac{\partial}{\partial p} \left(\frac{J}{\sigma} \right)$$

Term A

Term B

Term C

Term D
(6.23)

Physical meaning of (6.23) may be understood by the following simple form:

$$-\chi \propto -V_g \cdot \nabla (\zeta_g + f) + \frac{\partial}{\partial z} (-V_g \cdot \nabla T)$$

OR

$$-\chi \propto -V_g \cdot \nabla \zeta_g - \beta v_g + \frac{\partial}{\partial z} (-V_g \cdot \nabla T)$$

Term B: (1) Relative Vorticity Advection ($-V_g \cdot \nabla \zeta_g$)

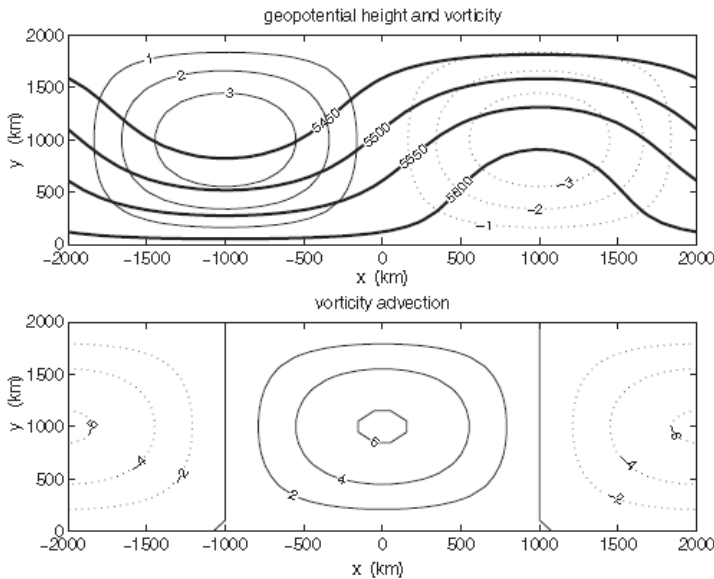
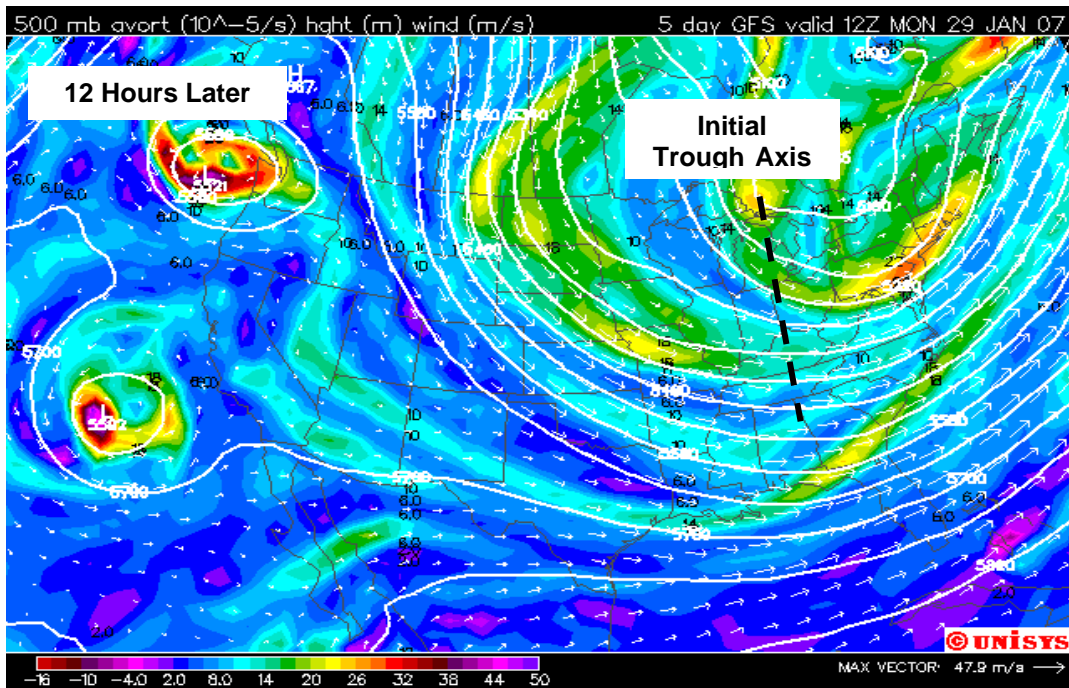
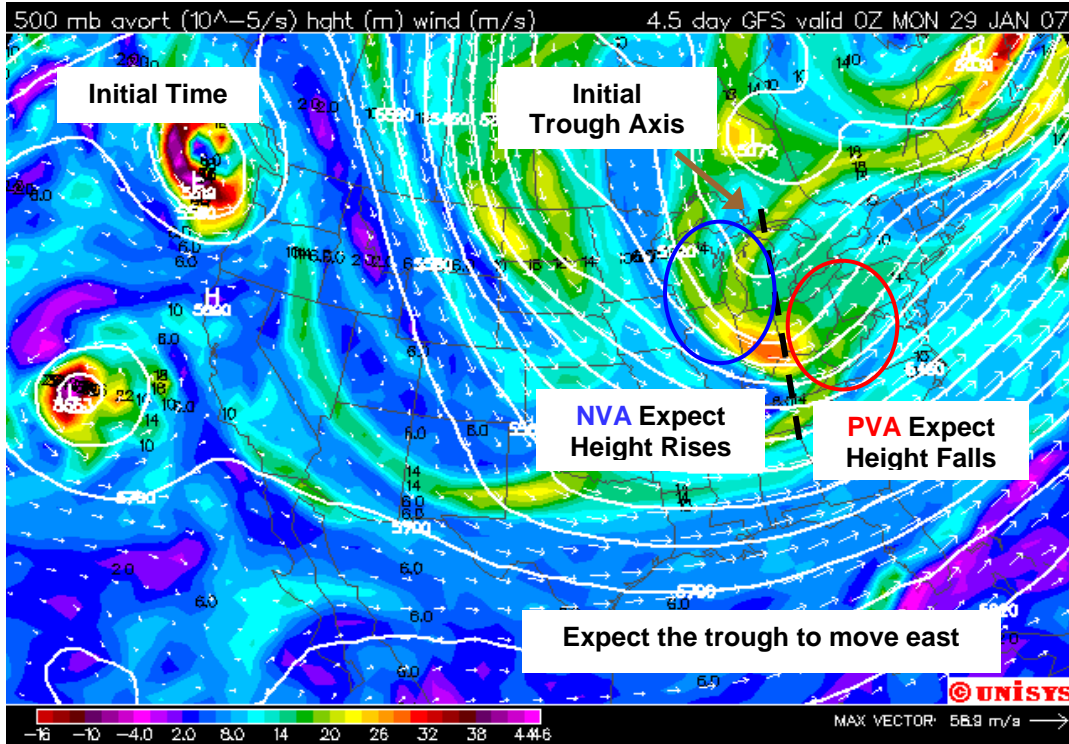


Fig. 6.8 (Top) Geopotential height in units of m, and relative vorticity in units of $10^{-5} s^{-1}$ for the sinusoidal disturbance of equation (6.20). Here $\Phi_0 = 5.5 \times 10^4 m^2 s^{-2}$, $f_0 = 10^{-4} s^{-1}$, $f_0 A = 800 m^2 s^{-2}$, $U = 10 m s^{-1}$, and $k = l = (\pi/2) \times 10^{-6} m^{-1}$. (Bottom) Advection of relative vorticity in units of $10^{-10} s^{-2}$ for the disturbance shown above.

(2) Planetary Vorticity Advection ($-\beta v_g$): Propagating

westward

Example of Relative Vorticity Advection



Term C: Differential Temperature Advection ($\frac{\partial}{\partial z}(-V_g \cdot \nabla T)$)

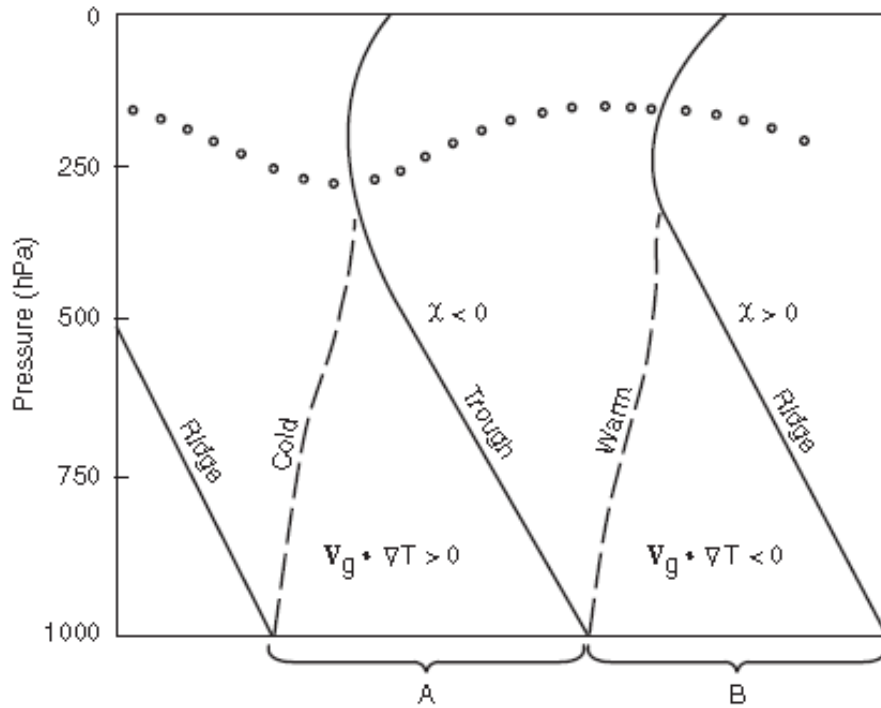


Fig. 69 East–west section through a developing synoptic disturbance showing the relationship of temperature advection to the upper level height tendencies. A and B designate, respectively, regions of cold advection and warm advection in the lower troposphere.

