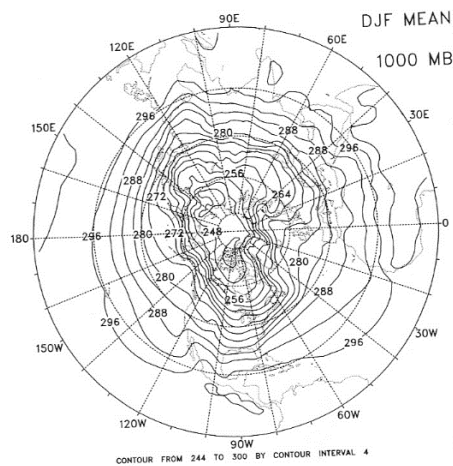


## Chapter 6 Introduction to General Circulation of the Atmosphere

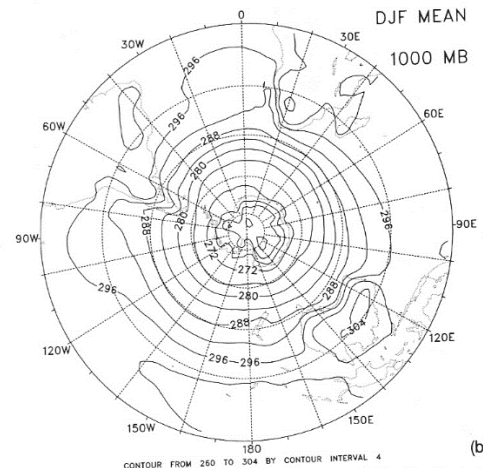
- Major features on sea-level synoptic charts

The major features include cyclones, fronts, and wave-like large-scale disturbances. The complicated patterns are mainly due to orography and land-sea contrasts.



(Bluestein 1993)

Northern Hemisphere



**Figure 1.44** Average temperature (K) at 1000 mb in (a) the Northern Hemisphere for the winter and (b) the Southern Hemisphere for the summer (December, January, and February) (from ECMWF data, 1979–1988; courtesy Kevin Trenberth and Amy Solomon, NCAR).

Southern Hemisphere

- **Major features on 500-mb charts**

The 500-mb charts are much smoother than sea-level synoptic charts. There are about 5 lows circling around the North Pole, which are associated with **baroclinic waves** (Fig. 3.1Hess).

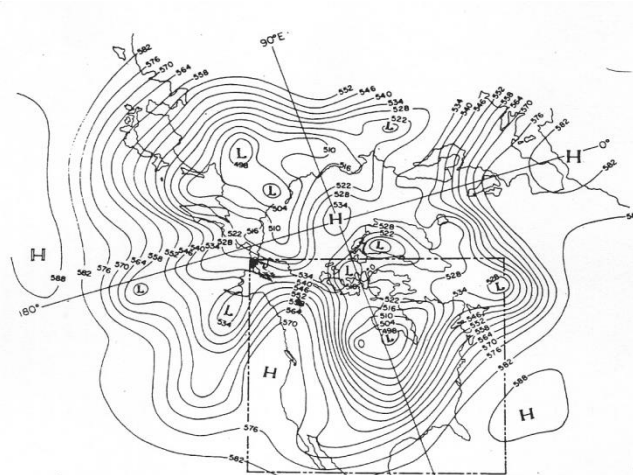


Fig. 3.1(Wallace & Hobbs): Distribution of geopotential height on the 500-mb surface at 00 UTC 20 November 1964. Labels on contours represent geopotential height (in m). The letters H and L denote centers of high and low geopotential height, respectively.

This type of **baroclinic waves** is often observed in the atmosphere, such as the 500-mb analysis chart shown below.

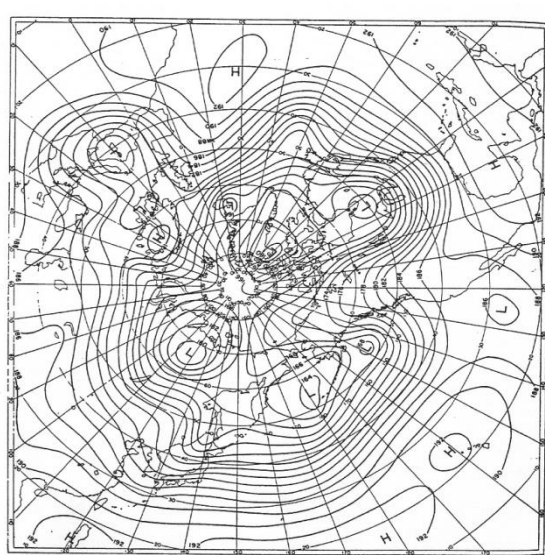
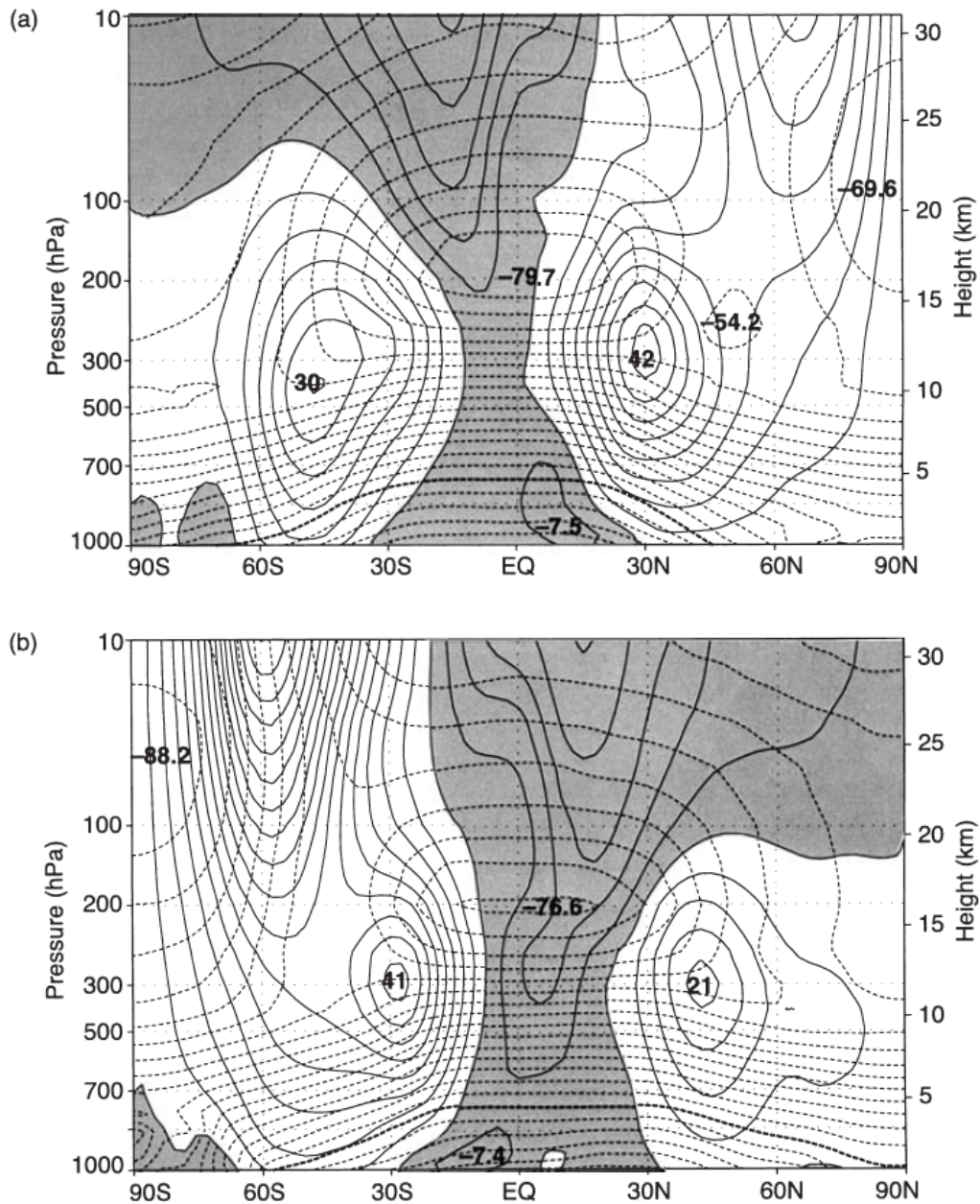
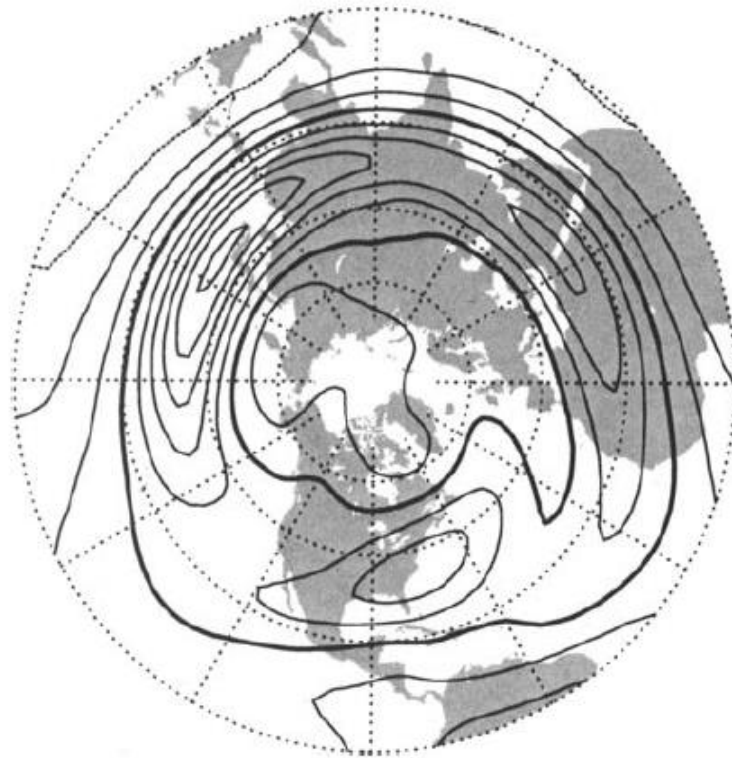


Fig. 6.1 (Holton): Meridional cross sections of longitudinally and time averaged zonal wind indicate strong seasonal variation.



**Fig. 6.1** Meridional cross sections of longitudinally and time-averaged zonal wind (solid contours, interval of  $\text{m s}^{-1}$ ) and temperature (dashed contours, interval of 5 K) for December–February (a) and June–August (b). Easterly winds are shaded and  $0^\circ\text{C}$  isotherm is darkened. Wind maxima shown in  $\text{m s}^{-1}$ , temperature minima shown in  $^\circ\text{C}$ . (Based on NCEP/NCAR reanalyses; after Wallace, 2003.)

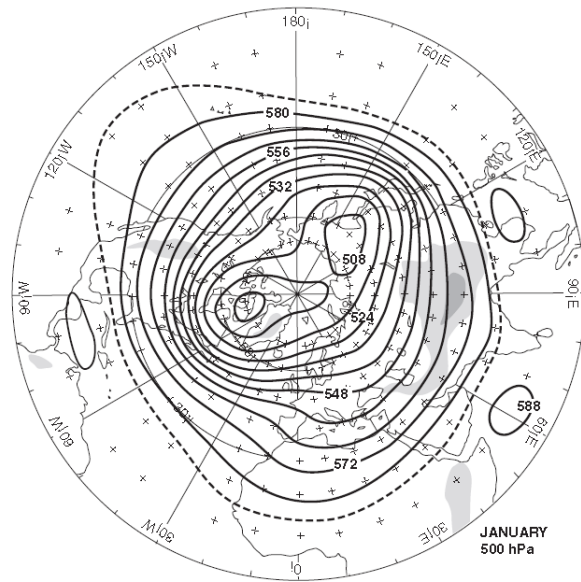
- Fig. 6.2 (Holton): Mean zonal wind (heavy contour, 20 m/s) at the 200-mb level for Dec-Feb averaged for 1958-1997 (Wallace 2003). Jet streams in midlatitudes; jet streams are not continuous longitudinally.



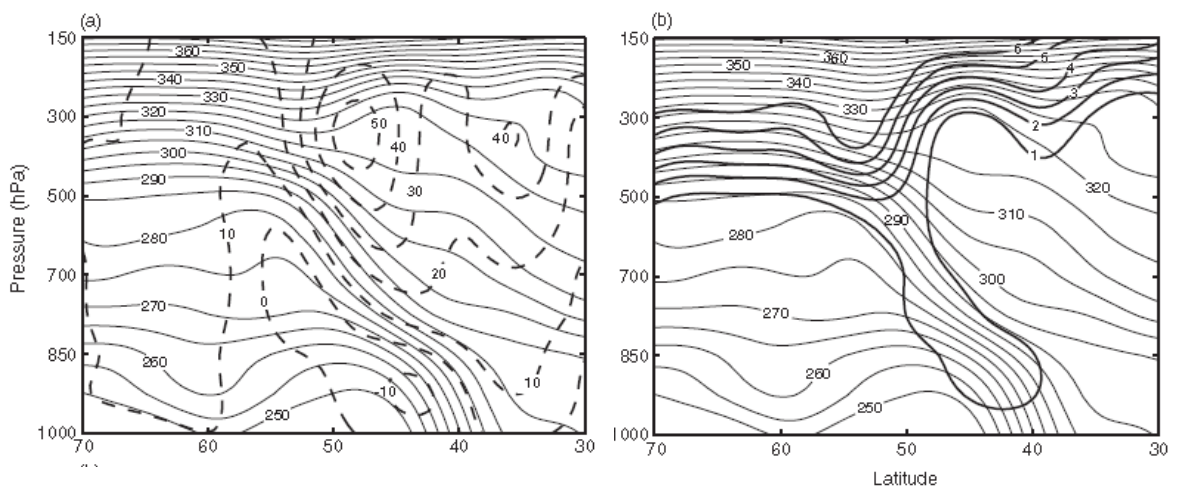
**Fig. 6.2** Mean zonal wind at the 200-hPa level for December–February averaged for years 1958–1997. Contour interval  $10 \text{ m s}^{-1}$  (heavy contour,  $20 \text{ m s}^{-1}$ ). (Based on NCEP/NCAR reanalyses; after Wallace, 2003.)



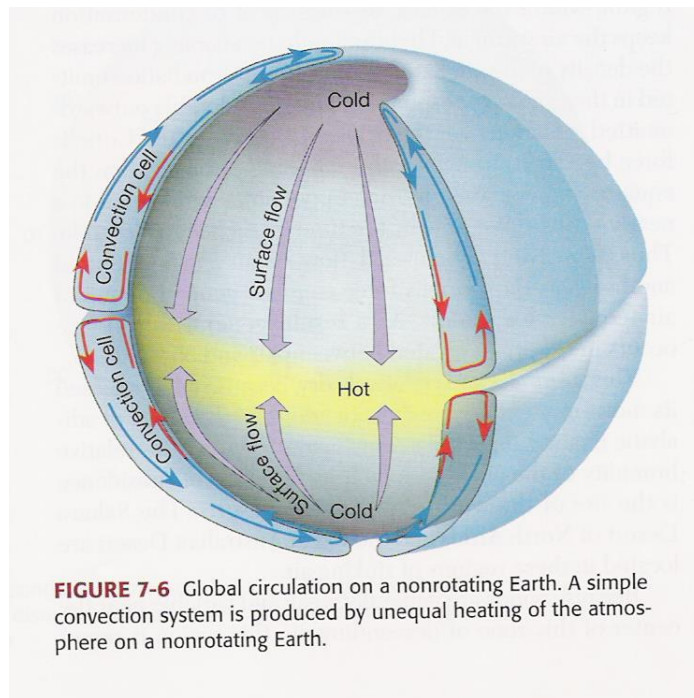
- Fig. 6.3 (Holton): Mean 500-mb height contours (in 10m) in January in the Northern Atmosphere (Palmen and Newton 1969).



- Fig. 6.4 (Holton): Latitude-height cross sections through a cold front: (1) strong baroclinicity near the front, (2) jet stream is collocated with the cold front due to thermal wind effect.



- General Circulation – The Hadley Cell Model (1735)



(Lutgens and Tarbuck 2004)

### The major problem with the Hadley Model!!!

There is net radiative loss of heat by the atmosphere associated with the upper-level flow toward poles. This cold air must subside and spread poleward and equatorward.

Thus the radiative consideration forces a breakdown of the single Hadley cell!

- Three-Cell Circulation Model (Ferrel 1859)

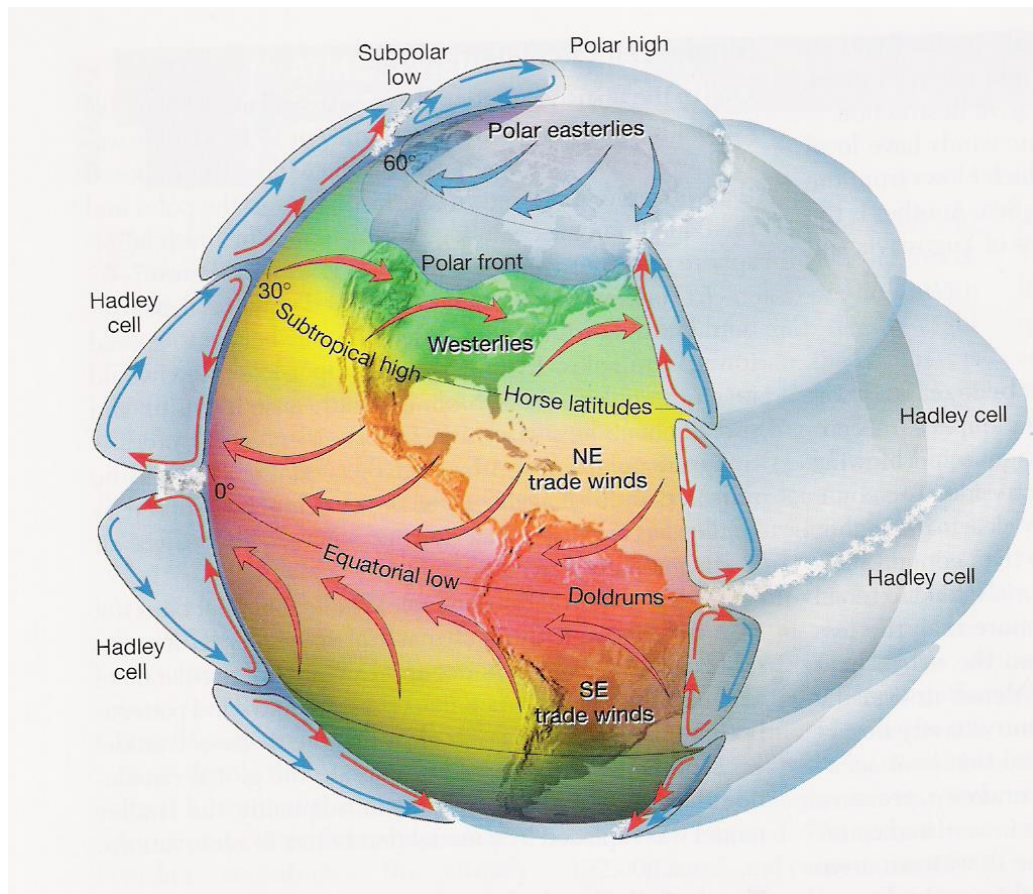
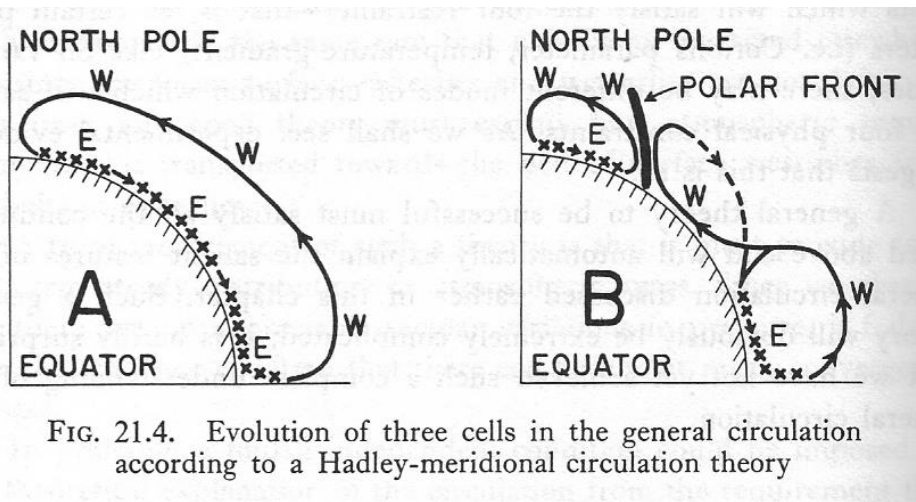


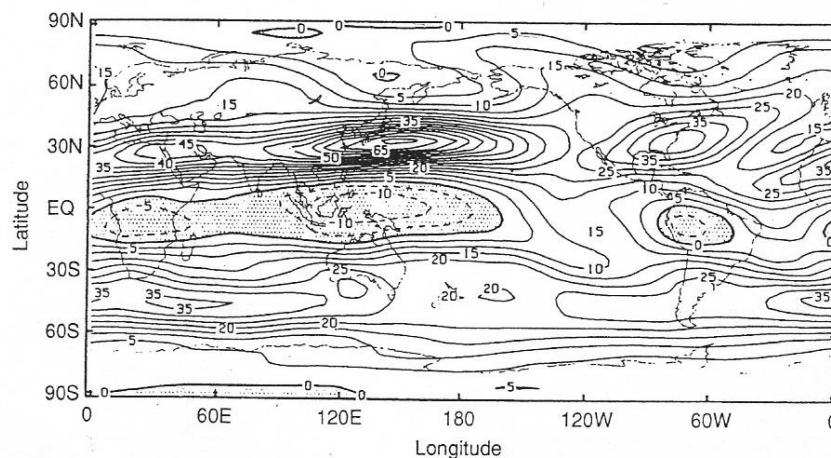
Figure 7.7 (Lutgens and Tarbuck 2004): Idealized global circulation proposed for the three-cell (Ferrel cell) circulation model of a rotating Earth.

- The 3-cell circulation was mainly driven by Earth's rotation and radiation on the upper branch of the Hadley cell.

- The problem with the three-cell model is that the theory fails to account for the fact that the westerlies of middle latitudes increase with height!
- Rossby has suggested that *large-scale turbulence* might link the upper westerlies of the Hadley cell and the polar cells to the upper troposphere of middle latitudes.



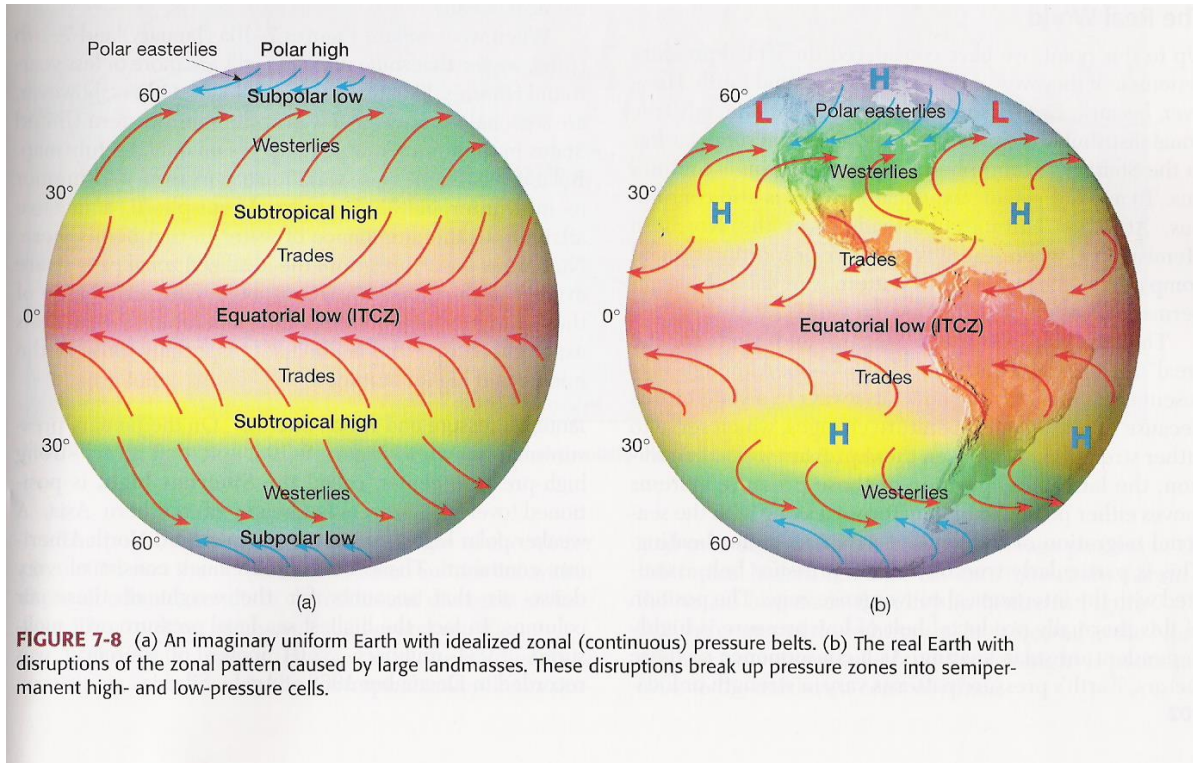
- The fact is that there exists a longitudinal variation of the flow, thus the jet stream is not continuous around the globe.



Latitude-longitude cross section of time-averaged zonal wind speed at 200 mb for DJF averaged for years 1980-1987. (After Schubert *et al.*, 1990.)



- Observed General Circulation of the Atmosphere



## Annulus Experiments

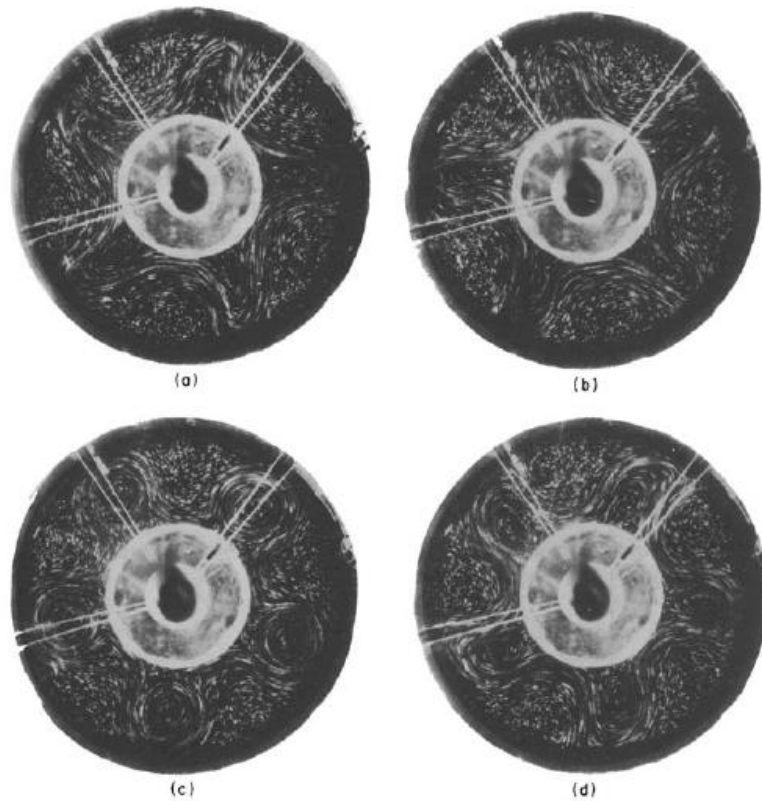
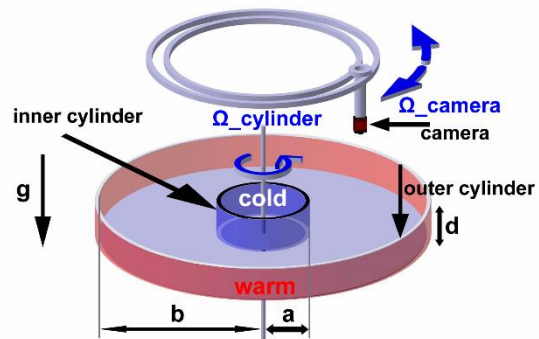


Fig. 10.20 (Holton): Time exposures showing the motion of surface tracer particles in a rotating annulus with heating at the rim and cooling at the core.

## Flow regimes found from annulus experiment:

### (1) Large Rossby number (low rotation rate) -- Hadley Regime

- Hadley cell circulation develops.
- The circulation is steady and symmetric.

### (2) Moderately large Rossby number (medium rotation rate)

- Meridional flow breaks down.
- Jet streams can be found near the upper surface.
- Very steady pattern of Rossby waves exists in westerlies.
- Number of waves evolves from 3, 4, to 5 as rotation rate increases.
- Motions are quasi-geostrophic.

### (3) Moderately small Rossby number (large rotation rate) -- Rossby Regime

- Flow is similar to regime 2
- Jets are shorter in length, not continuous
- Flow is markedly unsteady
- Five to six Rossby waves form
- Low-level cyclones and fronts appear
- Observed atmosphere features, such as cold anticyclones, are reproduced
- Troughs in the wave patterns tilted from NE to SW instead of lying along a meridian.

### (4) Small Rossby number (very high rotation rate)

- Physical circulation entities become smaller, which mimic convective Elements or Bernard cells.
- Cells are quite unsteady and short-lived.
- Narrow jets are winding in and out among the cells.